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# Dissolved organic matter in urban forestland soil and its interactions with typical heavy metals: a case of Daxing District, Beijing

Chen Zhao<sup>1</sup> · Shi-Jie Gao<sup>1,2</sup> · Lei Zhou<sup>1</sup> · Xiang Li<sup>1</sup> · Xi Chen<sup>1</sup> · Chong-Chen Wang<sup>1</sup>

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## Abstract

As an active substance, dissolved organic matter (DOM) acts a pivotal part in heavy metals (HMs) transportation from urban forestland soil to aquatic ecosystem. In this study, the soil samples from 35 individual subareas were scientifically collected with the aid of geographical information system (GIS) technology. UV-visible (UV-vis) and excitation-emission matrix (EEM)-related parameters suggested that the DOM in urban forestland soil mainly originated from terrestrial and microbial sources. Fluorescence quenching titration associated with parallel factor (PARAFAC) modeling was applied to quantify the complexation ability of four HMs (Cu, Cd, Pb, and Ni) and DOM in urban forestland soil. One fulvic-like (C1), two humic-like (C2 and C3), and one protein-like fluorophores (C4) were identified by EEM-PARAFAC modeling. Considerable differences in fluorescence quenching curves were observed between individual organic constituents and target HMs. Among the four HMs, addition of Cu(II) ions resulted in EEM spectra quenching of each PARAFAC-decomposed organic constituent. However, relatively strong fluorescence quenching phenomena were only detected in humic-like constituents (C2 and C3) with the titration of Pb(II) and Ni(II), which revealed that these types of organic constituent were predominantly responsible for Pb(II) and Ni(II) binding in urban forestland soil-derived DOM. Furthermore, considering the resistant nature of C2 and C3 constituents along with their significant quenching effects for the four target HMs, the concentrations of humic-like constituents in urban forestland soil may be a useful parameter to evaluate the potential risk of HMs immobilization and transformation.

**Keywords** Urban forestland soil · Dissolved organic matter · Characterization · Heavy metal · Quenching · EEM-PARAFAC

## Introduction

Urban forest acts a significant part in enhancing the quality of the natural environment. Well-managed urban forests can reduce stormwater runoff, relieve the pressure of the heat island, sequester CO<sub>2</sub>, and reduce air pollution (Yang et al. 2005). As the political, economic, and cultural center of China,

Beijing has experienced fast-speed urbanization and industrialization in the past few decades. The rapid development has put large amounts of pressure on its ecological environment. Heavy metal (HM) pollution is the critical issue, particularly in the urban forestland soil, which might originate from numerous anthropogenic activities like energy manufacturing (Al-Khashman and Shawabkeh 2006; Kumar et al. 2008), architecture industry (Ye et al. 2011), automobile exhaust (Ward et al. 1977; Wei and Yang 2010), waste disposal (Luo et al. 2011) and fuel combustion (Lind et al. 1999; Meza-Figueroa et al. 2007). The above activities discharge HMs into the air and they subsequently are accumulated into urban forestland soil as the HM-containing dry precipitation.

Dissolved organic matter (DOM) in soil ecosystems is a small but active part, which can interact with HMs to generate metal-organic complexes, thus exerting important influences on their morphology, distribution, biological toxicity, transportation, and transformation (Kalbitz and Wennrich 1998; Weng et al. 2002). In particular, the binding affinity of DOM with HMs can dramatically affect the soil quality, and also the

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✉ Chong-Chen Wang  
[chongchenwang@126.com](mailto:chongchenwang@126.com)

<sup>1</sup> Beijing Key Laboratory of Functional Materials for Building Structure and Environment Remediation, Beijing University of Civil Engineering and Architecture, Beijing 100044, China

<sup>2</sup> College of Forestry, Beijing Forestry University, Beijing 100083, China

mobilization and migration of HMs towards the aquatic environments (Ashworth and Alloway 2004; Refaey et al. 2017; Sun et al. 2017). Therefore, investigations on the interaction mechanism between HMs and DOM were significant to clarify the biochemical cycling of HMs in urban forestland soil environments.

The binding characteristics of HM-DOM were strongly related to the organic sources, compositions, and structures. For example, the humic-like and protein-like components extracted from municipal solid waste leachate were responsible for Cu(II) binding; whereas the protein-like component was the only fraction involved in Cd(II) complexation (Wu et al. 2012). Comparing with the Cu(II) ions, Zn(II) ions showed a lower affinity to DOM derived from lake sediments (Xu et al. 2013). Typically, phenolic hydroxyl and carboxylic ligands within DOM compositions were typically regarded as powerful and weak HMs binding sites, respectively (Croué et al. 2003; Hur and Lee 2011). Previous studies had verified that the capacities of DOM to bind HMs during their migration through the soil were one element that may prominently affect the quality of water environments. Because of the generation of resistant, aqueous complexes with HMs, such as Cu, Pb, and Hg, DOM can strengthen metal mobilization and transportation towards the groundwater (Ashworth and Alloway 2004; Dudal et al. 2005). Wu et al. (2003) suggested that some small organic acids (average molecular weight < 500 Da) in agricultural soil, such as acetic, oxalic, malic, and citric, can generate stable complexes with HMs. Yang et al. (2006) also found that massive small molecules of organic acids can affect the adsorption-desorption equilibrium of Pb in acidic clayey soil. Recently, researchers paid more attention to explore the effects of remediation materials (sewage sludge compost, food waste compost, biochars, etc.) on soil organic substances, revealing that they remarkably modified the chemical properties of DOM and subsequently influence the mobility and bioaccumulation of HMs (Beiyuan et al. 2018; Fang et al. 2016; Li et al. 2018a). However, up to now, few attentions were carried out to characterize the DOM in urban forestland soil and study its binding properties with co-existing HMs.

UV-visible (UV-vis) and excitation-emission matrix (EEM) spectroscopies are user-friendly, rapid, and non-destructive instruments that can supply useful information about the chemical compositions and molecular structures of DOM (Wang et al. 2018a; Zhao et al. 2015; Zhao et al. 2016; Zhao et al. 2018b). More specifically, EEM spectra combined with quenching analysis had been proven to be a reliable method to explore the interaction mechanism between HMs and organic substances (Chen et al. 2013; Yamashita and Jaffé 2008; Zhu et al. 2014). Introduction of parallel factor (PARAFAC) modeling to EEM spectroscopy can cut down the interference of overlapping peaks in original fluorescence spectra, obtain the binding parameters between metal ions and individual fluorophores in DOM, and provide much more

quantitative and specific information (Huang et al. 2018; Wang et al. 2018b; Wang et al. 2016; Wu et al. 2011; Wu et al. 2012). Ohno et al. (2008) verified that the binding characteristics between water soluble soil organic matter (WSOC) and trivalent metals (Fe and Al) might be well analyzed by EEM-PARAFAC modeling. Moreover, Xu et al. (2013) successfully estimated the binding characteristics of lake sediment-derived DOM and divalent metals (Zn and Cu) with the identical method.

In this study, field experiments were conducted in Daxing District, which is the largest region in southeast of Beijing, China. To make the subsequent analysis more representative and systematic, the sampling work was completed under the aid of geographical information system (GIS) technology. The objectives were (1) to use UV-vis and EEM spectroscopies to evaluate the chemical compositions and structures of DOM extracted from urban forestland soil in Daxing District; (2) to investigate the feasibility of EEM-PARAFAC and quenching method for application in quantification of the binding characteristics of HMs and DOM in forestland soil. Based on the above purpose, Cu, Cd, Pb, and Ni, which have comparatively high concentrations and biotoxicity in soil in Beijing (Chen et al. 2010), were selected as the representatives to discuss. Finally, this study may offer an in-depth view of the heterogeneous properties of HMs complexing to DOM existed in urban forestland soil at the molecular level and the influences of DOM on the environmental behaviors and the ecological risk of HMs in the soil ecosystems.

## Materials and methods

### Reagents and instruments

All chemical reagents were analytical grade from J&K Chemical Corporation (Beijing, China) and were used directly without any further purification. Each of the experimental solution was prepared with ultrapure water of resistance not less than  $18.2 \text{ M}\Omega \text{ cm}^{-1}$ . The contents of the four heavy metal titrants were standardized with the EDTA solution (Bai et al. 2008).

The pH values of the forestland soil samples were determined in a suspension (the ratio of water to dry soil was 2.5) utilizing a PB-10 pH meter (Sartorius, Germany). The total N and P of the soil samples were detected applying an AA3 continuous-flow analyzer (Seal Analytical Corporation, Germany). The HMs and dissolved organic carbon (DOC) contents in the extracted samples were measured by a Perkin-Elmer AA-PinAAcle 900 atomic absorption spectroscopy (Fermont, USA) and a Jena multi N/C 3100 analyzer (Jena, Germany), respectively. UV-vis and EEM spectra were determined on a Perkin-Elmer lambda 650S spectrophotometer (Fremont, USA) and a Hitachi F-7000 Fluorescence



spectrophotometer (Tokyo, Japan), respectively. Standard quinine sulfate units (QSU) were selected to compare the relative concentrations of fluorophores, in which 1 QSU was equivalent to 32.2 intensity unit. The primary scatters (Rayleigh Raman scatters) were removed by setting zero to the EEM spectra in the two triangle regions ( $E_x \leq 0.5 E_m + 5 \text{ nm}$  and  $\geq E_m - 20 \text{ nm}$ ) (Stedmon and Bro 2008), and by subtracting the blank spectra (ultrapure water in this study) (Santos et al. 2009), respectively.

### Soil sampling and pretreatment

The sampling area of this study was concentrated in the Daxing District, Beijing. The total sampling area was approximately  $135.28 \text{ km}^2$ . The whole area was divided into 35 typical subareas based on GIS technology, which guaranteed the accuracy of the sampling locations (Fig. 1). Average 5 samples were collected at each sampling subarea using Soil Core Samplers (AMS samplers, USA).

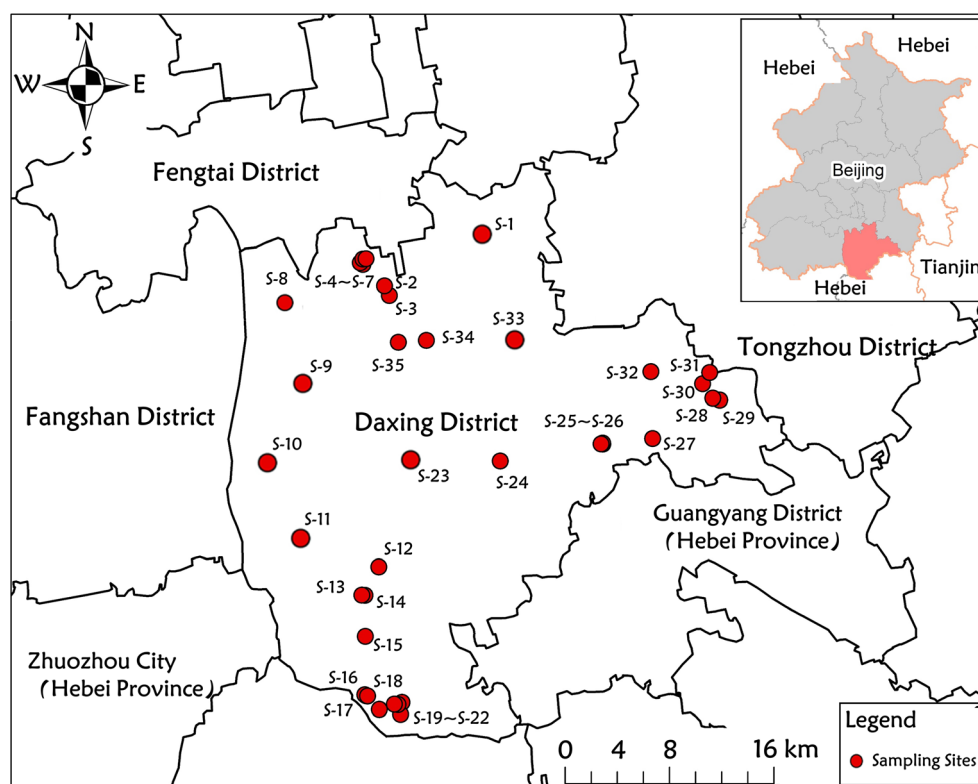
In order to facilitate the transportation and follow-up analysis, the collected soil samples were reserved in polyethylene bags. The soil samples were placed in an oven at  $50^\circ\text{C}$  and dried for 3 days (Del Galdo et al. 2003, Lee et al. 2006; Six et al. 2002). After that, they were filtered through a 2.0-mm sieve to eliminate stones, debris, and other interfering substances. Approximately 20 g soil samples were ground in a mechanical agate grinder until fine particles ( $< 200 \mu\text{m}$ ) were

obtained. Remarkably, DOM constituents were extracted by mixing one part of well-prepared soil sample with four parts of ultrapure water and continuously shaking them for 24 h. The extracts were centrifuged for 10 min at 8000 rpm at  $4^\circ\text{C}$  and passed through a  $0.45\text{-}\mu\text{m}$  hydrophilic PVDF Millipore membrane (Nierop et al. 2002; Yu et al. 2012), and the physico-chemical parameters of DOM were shown in Table 1.

### Fluorescence quenching titration

Prior to fluorescence titration, DOM in extracted solutions with absorbance greater than 0.02 at 350 nm will be diluted with ultrapure water, and the ionic strength of the DOM solutions were controlled by  $0.1 \text{ mol L}^{-1} \text{ KClO}_4$  to avoid the inner filter effects (Ohno 2002). The pH values of the solutions were adjusted to the expected values (pH = 4.0 for  $\text{Cu}(\text{NO}_3)_2$ , pH = 4.5 for  $\text{Pb}(\text{NO}_3)_2$  and  $\text{Ni}(\text{NO}_3)_2$ , pH = 6.0 for  $\text{Cd}(\text{NO}_3)_2$ ) by adding  $0.1 \text{ mol L}^{-1} \text{ HNO}_3$  or  $\text{NaOH}$  solutions, and the total addition should not exceed  $200 \mu\text{L}$ . Both aliquots of 25 mL diluted DOM extract and  $0.02 \text{ mol L}^{-1} \text{ HM}$  standard solutions were put into 40-mL brown sealed vials. The HM concentrations in the ultimate solutions remained in the range of  $0\text{--}600 \mu\text{mol L}^{-1}$  with the content interval of  $40 \mu\text{mol L}^{-1}$ . All mixed solutions of DOM and HMs were shaken for 24 h without light irradiation at ambient temperature to guarantee equilibrium of complexing reactions (Zhao et al. 2018a). In addition, the same volume of DOM solution with the ionic

**Fig. 1** Schematic map of the target area and the sampling sites



**Table 1** Background chemical characteristics of the collected urban forestland soil samples ( $n = 35$ )

Values	pH	TN g kg <sup>-1</sup>	TP	TOC mg L <sup>-1</sup>	Cu μg L <sup>-1</sup>	Cd	Pb	Ni
Range	6.57–7.80	0.02–0.11	0.05–0.51	3.61–16.85	7.11–27.25	0.51–5.28	0.21–2.88	0.50–14.20
Mean	7.23	0.06	0.21	9.50	13.35	1.86	0.94	5.32
Median	7.43	0.06	0.19	8.69	15.12	1.27	0.83	6.21
RSD (%)	7.12	49.45	70.82	35.56	42.67	37.28	31.07	45.21

strength of 0.1 mol L<sup>-1</sup> KClO<sub>4</sub> was selected as the blank to accomplish EEM fluorescence measurement.

### PARAFAC modeling

PARAFAC model can systematically decompose and extract individual organic substances from EEM spectral dataset. It also offers the relative abundance of individual organic constituent in each sample (Ohno et al. 2008; Stedmon and Bro 2008). All EEM spectra data in this study were compiled into three-way arrays of  $X$  with  $I \times J \times K$ , where  $I$ ,  $J$ ,  $K$  represent the amount of samples, emission, and excitation wavelengths, respectively. The modeling can be described using Eq. (1):

$$X_{ijk} = \sum_{f=1}^F a_{if} b_{jf} c_{kf} + \varepsilon_{ijk} \quad i = 1, \dots, I, j = 1, \dots, J, k = 1, \dots, K \quad (1)$$

where  $F$  is the amount of fluorophores,  $X_{ijk}$  defines the fluorescent peak of sample  $i$  at given emission  $j$ /excitation  $k$  wavelength;  $a_{if}$  is proportional to the content of the  $f_{th}$  organic constituent in the  $i_{th}$  sample;  $b_{jf}$  and  $c_{kf}$  are assessments of the emission and excitation spectra, respectively, for the  $f_{th}$  organic constituent. Finally,  $\varepsilon_{ijk}$  defines the residual matrix to minimize the sum of squared residuals by changing least squares algorithm (Stedmon et al. 2003).

EEM-PARAFAC analysis (99 samples  $\times$  55 emissions  $\times$  55 excitations) was completed applying the DOMFluor toolbox in MATLAB 7.0 software (Mathworks, Natick, MA). The relative concentrations of individual fluorophores were described by maximum fluorescence intensities ( $F_{max}$ ). Split-half analysis was used to prove the correctness of the decomposed data from PARAFAC modeling.

### Complexing reaction modeling

In this research, the Ryan-Weber model was used to quantify the binding parameters between HMs and individual organic constituents decomposed from EEM-PARAFAC modeling, which can provide more accurate and detailed information than the traditional peak maxima method (Ryan and Weber 1982a). The principle of this model assumes that organic

ligands complex with HMs at a molar ratio of 1:1, which can be expressed as Eq. (2):

$$I = I_0 + (I_{ML} - I_0) \left( \frac{1}{2K_M C_L} \right) \times \left( 1 + K_M C_L + K_M C_M - \sqrt{(1 + K_M C_L + K_M C_M)^2 - 4K_M^2 C_L C_M} \right) \quad (2)$$

where  $I_0$  and  $I$  define the fluorescence intensity without the titration and at the HM concentration of  $C_M$ , respectively;  $I_{ML}$  is the limiting fluorescence intensity;  $K_M$  and  $C_L$  are the conditional stability constant and binding capacity, respectively (Yamashita and Jaffé 2008). The fraction of the binding fluorescent components ( $f$ ) was obtained by Eq. (3):

$$f = \frac{(I_0 - I_{ML})}{I_0} \times 100 \quad (3)$$

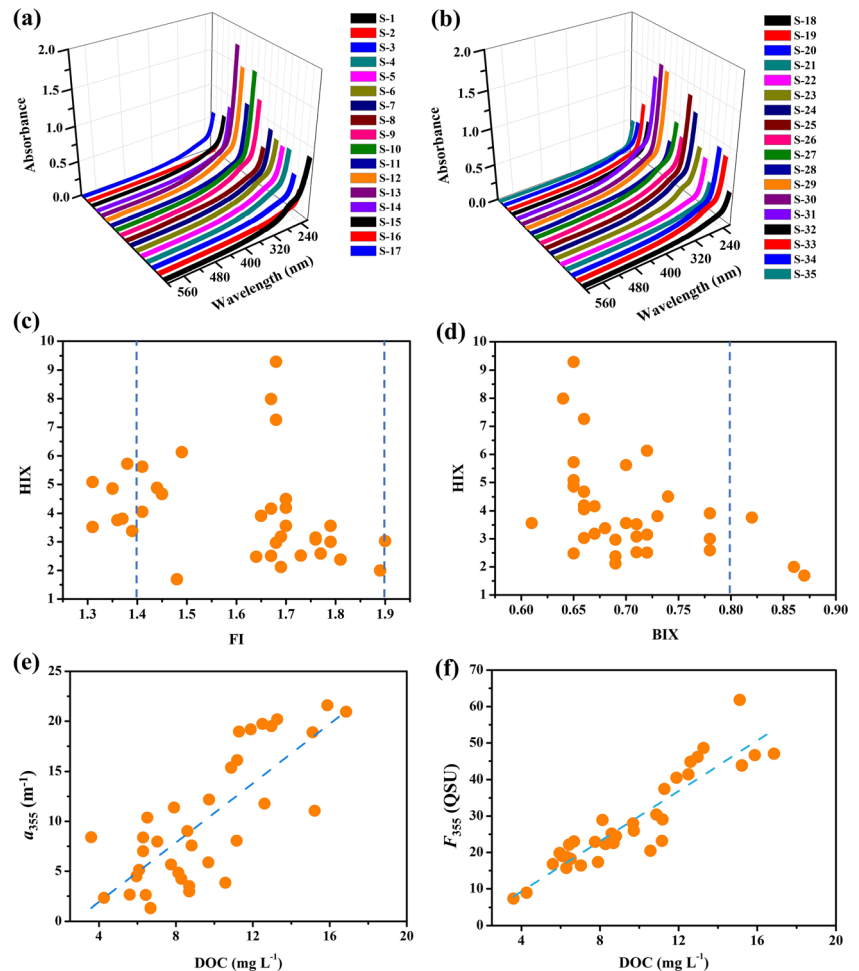
## Results and discussion

### Characterization parameters calculated by UV-vis and EEM spectral data

The UV-vis spectra of DOM extracted from urban forestland soil samples at different sampling sites were illustrated in Fig. 2a, b, which increased approximately exponentially with decreasing wavelength. Similar to previous researches (Hong et al. 2014; Jiang et al. 2017), no prominent absorption peaks were found in the UV-vis spectra, only inconspicuous shoulder peaks were detected in the range of 250–300 nm, which could be attributed to presence of conjugated molecular structures ( $\pi \rightarrow \pi^*$  electron transitions) in DOM. To gain more information from these spectra about DOM of forestland soil samples, the spectral slope coefficient ( $S_{275-295}$ ,  $\mu\text{m}^{-1}$ ), specific UV absorbance at 254 nm ( $SUVA_{254}$ ,  $\text{L mg}^{-1} \text{C}^{-1} \text{m}^{-1}$ ), the ratio of UV absorbance at 254 and 203 nm ( $A_{253/203}$ ), and absorption coefficient at 355 nm ( $a_{355}$ ,  $\text{m}^{-1}$ ) were calculated in this study (the corresponding values were shown in Table 2 and their computational details were revealed in Electronic Supplementary Information).

It is apparent that  $S$  is one of the important parameters to characterize DOM. However, the validity of its application

**Fig. 2** **a, b** UV-vis spectra of DOM extracted from urban forestry soil samples at different sampling sites. **c, d** Relationships of FI vs HIX and BIX vs HIX. **e, f** Correlations plots between DOC concentrations and CDOM ( $r^2 = 0.59$ ) and FDOM ( $r^2 = 0.83$ )



mainly depends on the calculation range of wavelength. For example, Stedmon et al. (Stedmon and Markager 2001) obtained  $S$  values by calculating the spectral data from 300 to 650 nm to characterize the DOM in the Greenland Sea. Santos et al. (2012) fitted  $S$  values within the spectra 240–400 nm to estimate the average molecular weight of DOM in rainwater. In this study, wavelengths between 275 to 295 nm were chosen to calculate the  $S$  values due to relatively high absorbance levels in this wavelength range. Furthermore, the  $S_{275-295}$  has been proven to negatively correlate with the molecular weight of organic substances (Helms et al. 2008). The  $S_{275-295}$  values determined for the DOM extracted from urban forestland soil

samples were in the range of 12.11–16.20  $\mu\text{m}^{-1}$  with the median of 13.66  $\mu\text{m}^{-1}$ , which were higher than those reported for the other soil organic substances (as shown in Table 3), suggesting that the DOM in the urban forestry soil collected from the above sampling sites had smaller molecular weight. In addition, Spencer et al. (2012) found that the  $S_{275-295}$  values could negatively correlate with the hydrophobic organic components, so the hydrophobicity of DOM in the urban forestland soil was weaker than the other soil ecosystems.

$SUVA_{254}$  has been diffusely applied to evaluate the aromaticity of DOM; as illustrated in Table 2, the  $SUVA_{254}$  values of DOM were in the range of 2.83–6.81  $\text{L mg C}^{-1} \text{m}^{-1}$  with the

**Table 2** DOM optical properties of the collected urban forestland soil samples ( $n = 35$ )

Values	$S_{275-295}$ $\mu\text{m}^{-1}$	$SUVA_{254}$ $\text{L mg C}^{-1} \text{m}^{-1}$	$A_{253/203}$	$a_{355}$ $\text{m}^{-1}$	$F_{355}$ QSU	FI	HIX	BIX
Min	12.11	2.83	0.09	1.31	7.33	1.31	1.69	0.61
Max	16.20	6.81	0.27	21.60	61.80	1.90	9.29	0.87
Mean	13.79	4.69	0.19	10.09	28.23	1.72	4.02	0.70
Median	13.66	4.88	0.18	8.38	23.15	1.70	3.56	0.69

**Table 3** Summary of UV absorption related parameters of DOM in typical soil environments

Samples	$S_{275-295}$ $\mu\text{m}^{-1}$	References	Samples	$SUVA_{254}$ $\text{L mg}^{-1} \text{C}^{-1} \text{m}^{-1}$	References
Urban forestland soil	12.11–16.20	This study	Urban forestland soil	2.83–6.81	This study
Soil leachate	11.20	(Yang et al. 2015)	Fluvic hypercalcaric cambisol	1.50–3.00	(Hassouna et al. 2012)
Boreal forest soil	11.01–12.84	(O'Donnell et al. 2016)	Grassland soil	1.36	(Stumpe and Marschner 2010)
Permafrost	10.80–15.20	(Gao et al. 2018)	Boreal forest soil	< 4.50	(O'Donnell et al. 2016)
Farmland soil	13.00–16.00	(Jiang et al. 2012)	Saline-alkali soil	< 2.83	(Li et al. 2017)

median value of  $4.88 \text{ L mg C}^{-1} \text{m}^{-1}$ . Being compared to the previous studies about DOM in the other types of soil environments (as shown in Table 3), the corresponding higher  $SUVA_{254}$  values demonstrated that much more aromatic substances present in the urban forestland soil samples. The  $A_{253/203}$  can offer further information about substituted degree of aromatic DOM structures. Unsubstituted aromatic ring structures typically exhibit a relatively lower value of  $A_{253/203}$ , and an increase of  $A_{253/203}$  value suggests the organic components contain more polar functional groups (i.e., carboxylic, carbonyl, hydroxyl, and ester groups) on aromatic rings (Korshin et al. 1997). The median value of  $A_{253/203}$  (0.18) was slightly higher than commercial soil fulvic acids Waskish Peat (0.16) and Pony Lake (0.11) from the International Humic Substance Society (IHSS) (Mignone et al. 2012), implying that DOM in urban forestry soil contained more substitute functional groups and aliphatic chain structures.

In order to further reveal the DOM characteristics, three fluorescent indices, namely the fluorescence index (FI), the humification index (HIX), and the biological index (BIX), were calculated in this study (the corresponding computational details were shown in Electronic Supplementary Information). Among the above three indices, FI can be applied to identify DOM sources, such as autochthonous ( $> 1.9$ ) and allochthonous sources ( $< 1.4$ ) (Huguet et al. 2009; Mcknight et al. 2001). As illustrated in Fig. 2e, the FI values of most urban forestland soil samples were between 1.4 and 1.9, suggesting that the DOM might originate from both terrestrial and microbially derived organic substances. Additionally, some of soil samples that exhibited their FI values were less than 1.4, which may be ascribed to great anthropogenic activities. HIX values generally increase as biomass decomposition (Hunt and Ohno 2007), while low HIX values ( $< 10$ ) indicate that the non-humified structures were dominated in organic components (Huguet et al. 2009). Table 2 showed that the DOM samples from different sampling sites had low HIX values, ranging from 1.69 to 9.29 (Table 2), suggesting that the humification degree of DOM in urban forestland soil was relatively lower, which could be interpreted by two potential reasons: (1) the afforestation time was short (generally less than 10 years), the DOM mainly originated from incomplete decomposition of plant litter, soil humus or root exudates; (2) a massive

microorganic production caused by human activities (protein and sugar) might lead to “priming effect,” which could stimulate the microbial community to degrade humified organic substances (Bianchi 2011). BIX is a proxy of autochthonous microbially produced DOM inputs (Huguet et al. 2009). The BIX with a high value ( $> 0.8$ ) represents prominent autochthonous sources of freshly generated organics (Huguet et al. 2009; Huguet et al. 2010). All BIX values in this research varied from 0.61 to 0.87, with the median value of 0.69. This result indicated that most of organic substances in urban forestry soil were terrestrial sources.

The parameters  $a_{355}$  and  $F_{355}$  were considered as two quantitative indicators of chromophoric DOM (CDOM) and fluorescence DOM (FDOM), respectively (Chen et al. 2004; Guo et al. 2007; Li and Hur 2017). The  $a_{355}$  varied from 1.31 to  $21.60 \text{ m}^{-1}$ , and the  $F_{355}$  varied from 7.33 to  $61.80 \text{ QSU}$ . Both  $a_{355}$  and  $F_{355}$  correlated positively with DOC concentrations (as illustrated in Fig. 2e, f), revealing that CDOM and FDOM were significant and substantial constituents to the dissolved organic pool in urban forestland soil (Jiang et al. 2017; Zhao et al. 2015). Moreover,  $a_{355}$  and  $F_{355}$  could be applied as indicators for DOC concentrations via in situ detection in this type of soil ecosystem.

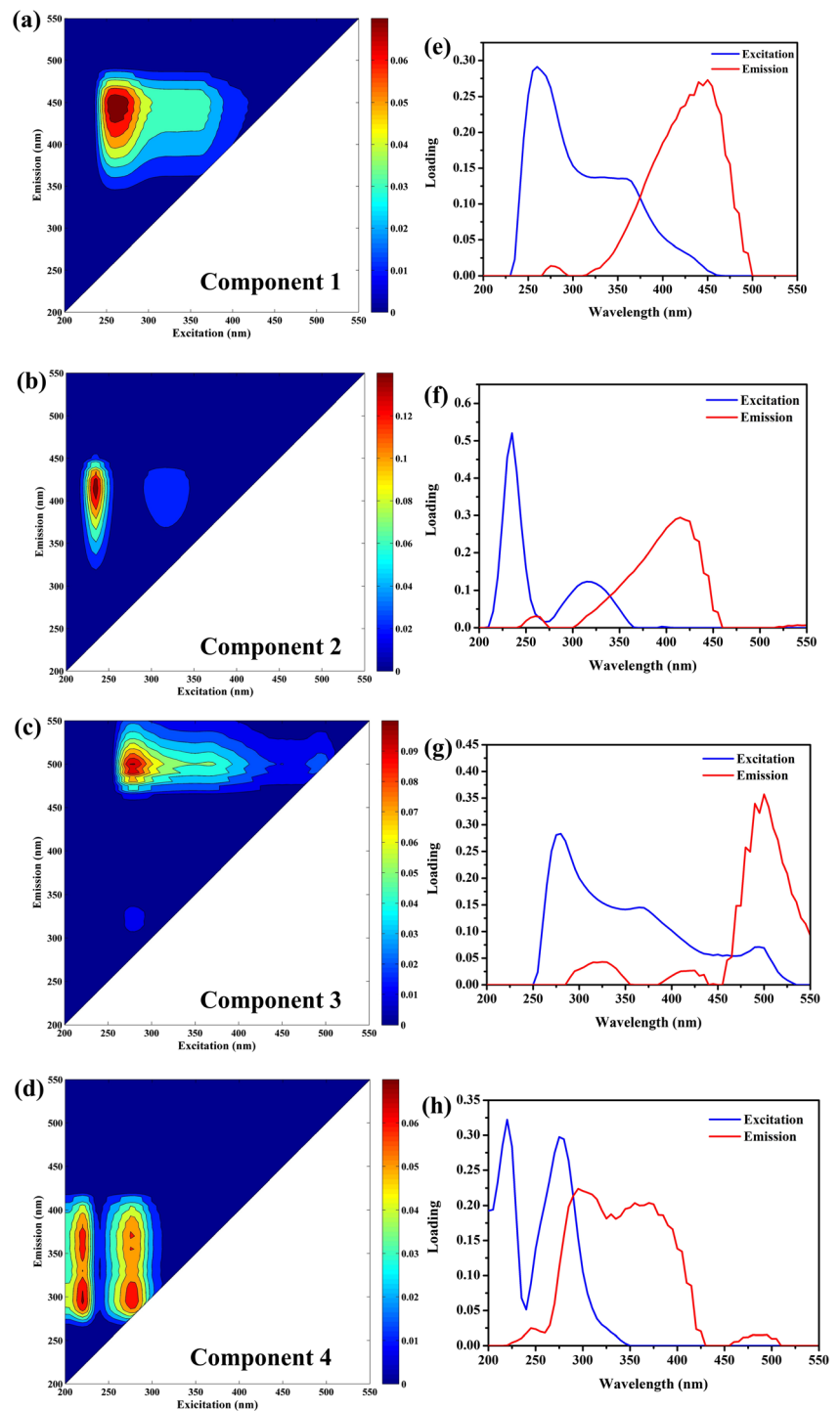
### PARAFAC analysis of urban forestland soil DOM

All the fluorescence spectral data of the urban forestland soil samples titrated with four HMs were decomposed by the PARAFAC modeling (Fig. 3). Four organic constituents were distinguished by the PARAFAC from the urban forestland soil DOM, including one fulvic-like constituent (C1), two humic-like constituents (C2 and C3) and one protein-like constituent (C4).

C1 presented one peak at Ex/Em wavelength of 260/450 nm (Fig. 1a, e). This spectral feature was considered as the fulvic-like fluorescence, consistent with previous researches (Liu et al. 2017; Mounier et al. 2011; Zhou et al. 2014). C2 showed a dominating and a secondary peaks at 230/423 nm (Ex/Em) and 315/423 (Ex/Em), respectively. This constituent was similar to the humic acid fluorophores and termed peak C by Coble (1996). C1 and C2 are mainly derived from terrestrial organic precursors like soil extracts, forested streams, wetlands, and deciduous leaves (Fellman et al. 2008; Fellman et al. 2009;



**Fig. 3** EEM fluorophores decomposed by PARAFAC model from the urban forestland soil DOM



Stedmon and Markager 2005; Stedmon et al. 2003). Wet precipitation generally leads to an increase in C1 and C2, which can be ascribed to enhance land-based inputs from stormwater runoff (Stedmon and Markager 2005). It is noteworthy that Ishii and co-workers (Ishii and Boyer 2012) had classified the C1 and C2 constituents as the UVC humic-like constituents based on their peak excitation wavelengths and

representative spectra. In view of the limitation of UVC light sources, C1 and C2 identified in this study were expected to be relatively resistant to photodegradation. In addition, Ohno et al. (2010) found that the molecular weight of C1 and C2 constituents in soil ecosystems was much more lower than the other co-existing organic constituents, and the average molecular weight was generally less than 665 Da.

The EEM spectral characteristics of C3 exhibited its peak at 275/500 nm (Ex/Em), which were an integration of terrestrial humic-like fluorescence peak A and peak C (Coble 1996). Lee et al. (2015) reported that fluorophores at short wavelengths were related to the existence of low aromatic organic constituents, while fluorescent organic constituent at long wavelengths were associated with high aromatic polycondensation structural organics. Being compared to C1 and C2, the peak location of C3 was shifted to longer wavelengths. Thus, among the three identified humic-like constituents, C3 organic constituent was probably associated with the highly condensed structures and larger molecular weight, indicating a relatively more refractory and mature nature.

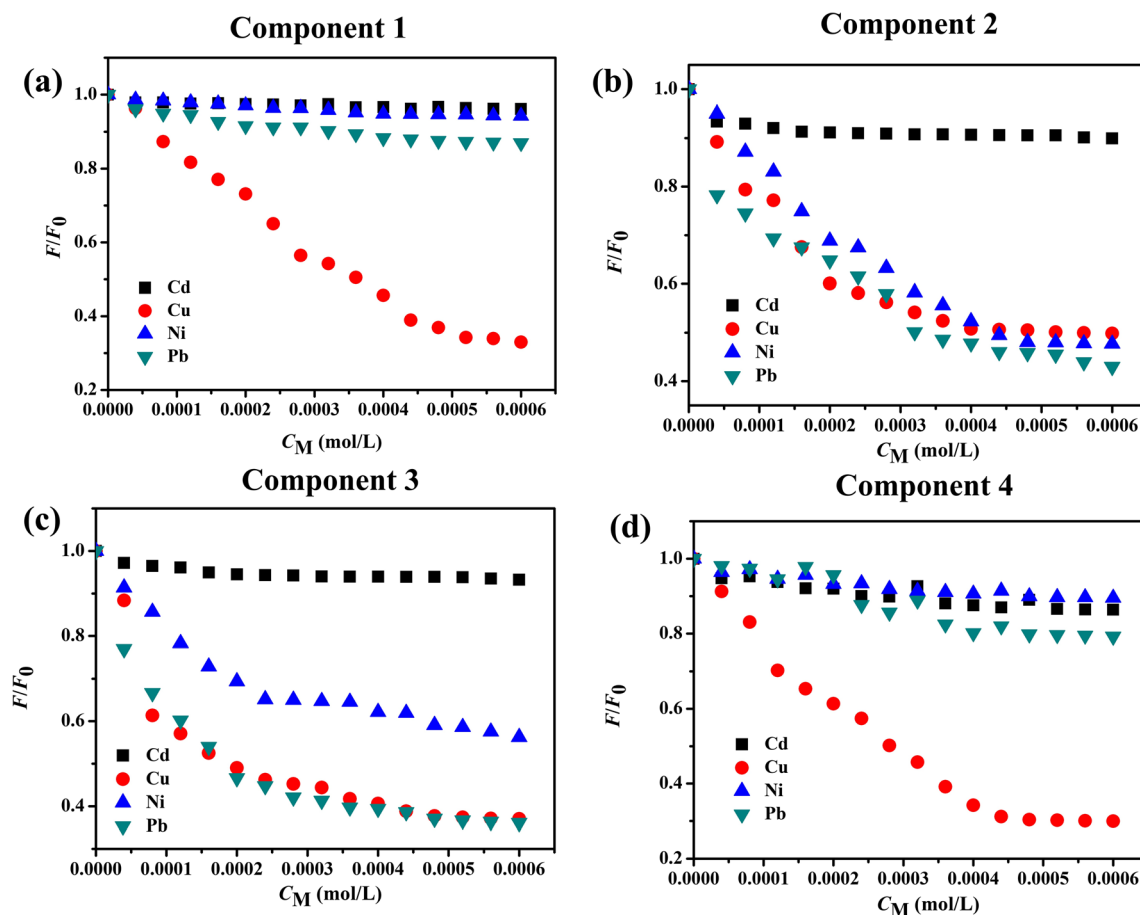
C4 had its maximum at 225/300, 225/375, 275/300, and 275/375 nm (Ex/Em), respectively. The bottom two peaks (225/300 and 275/300 nm) in the EEM spectrum (Fig. 3d) represented tyrosine-like peak B constituents (Coble 1996; Li et al. 2018b; Zhou et al. 2016). The other two peaks (225/375 and 275/375 nm) were attributed to tryptophan-like peak T constituents (Coble 1996; Li et al. 2018b), which were frequently detected in the aquatic environments affected by anthropogenic input (Hong et al. 2012; Phong and Hur 2015).

The results suggested that the DOM in urban forestland soil samples contained almost all types of dissolved amino acids (tyrosine- and tryptophan-like constituents), and its chemical composition was susceptible to human activities.

PARAFAC modeling offered excess quantitative information depicting the percentage composition of the fluorophores in the urban forestland soil samples. By calculating the  $F_{\max}$  values, the DOM displayed the higher abundance of C1 (26.69–47.59%) and C2 (29.44–41.12%), followed by C3 (12.62–23.18%) and C4 (5.80–19.94%). The total  $F_{\max}$  value of C2 and C3 occupied a predominant proportion in the total organic constituents, indicating that the humic-like organic constituents were the primary sources of the urban forestland soil in Daxing district.

### Interactions of PARAFAC-derived constituents with the HMs

The EEM quenching curves of individual constituent with the titration of Cu(II), Cd(II), Ni(II), and Pb(II) were illustrated in Fig. 4. For all fluorophores, a distinct quenching phenomenon was detected for Cu(II), while negligible quenching



**Fig. 4** Variation in the relative fluorescent intensity of individual PARAFAC-derived fluorophores with the addition of Cu(II), Cd(II), Pb(II), and Ni(II) solutions

phenomenon was detected for Cd(II), which were extremely similar to the widespread DOM in aquatic environments (Yamashita and Jaffé 2008), municipal solid waste leachate DOM (Wu et al. 2011), and commercial humic acid (Divya et al. 2009). Interestingly, the titration of Ni(II) and Pb(II) ions had comparatively strong quenching effects on C2 and C3, whereas the quenching effects on C1 and C4 was relatively weak. Weng et al. (2002) applied WHAM model to predict the affinity of sandy soil DOM for heavy metals, following the order:  $\text{Co} < \text{Ni} < \text{Cd} \sim \text{Zn} < \text{Pb} \sim \text{Cu}$ . The contradictory results of this study can be attributed to the differences in the functional group, molecular weight, and chemical composition of DOM in urban forestland soil. Previous studies had demonstrated that ligands such as phenols ( $-\text{OH}$ ), amines ( $-\text{NH}_2$ ), and carboxyl ( $-\text{COOH}$ ) may act a significant role in the complexing reaction (Struthers et al. 2008), and HMs with high binding strength were distributed more in the larger DOM fractions (Wu et al. 2004). Therefore, C2 and C3 contained much more active groups that can coordinate with HMs. Moreover, Jordan et al. (1997) concluded that peat-derived humic acid had higher affinity than fulvic acid for Pb, which further verified the accuracy of the PARAFAC model for the classification of four organic constituents in urban forestland soil samples.

With respect to protein-like fluorophores, the complex reaction with HMs might enhance or quench their fluorescent intensities (Wu et al. 2011; Yamashita and Jaffé 2008). Recently, EEM quenching method is diffusely applied to quantify the binding properties between humic-like substances and HMs and has seldom been used to determine the HMs binding parameters of given protein-like or amino acid molecules. Being compared to those of C1–C3, obvious fluctuations were examined in the quenching curves of C4, especially for Cd(II), Ni(II), and Pb(II) ions. This phenomenon can be proved by previous researches that humic-like and protein-like constituents can potentially generate supramolecular assemblies by means of  $\pi$ - $\pi$  interactions, van der Waals forces, and the other weak dispersive forces (Peuravuori and Pihlaja 2004; Romera-Castillo et al. 2014; Wang et al. 2015). Hence, the enhancement of fluorescent intensities after the complexation of protein-like constituents with HMs was largely attributed to metal-induced steric hindrance, which resulted in the reduction of interaction

between protein-like and humic-like constituents (Yamashita and Jaffé 2008). In this research, the fluorescent intensities of C4 were strongly quenched by Cu(II) ions without any fluctuation, indicating that fluorescence quenching method can be utilized to assess the binding parameters of protein-like constituents in urban forestland soil and Cu(II) ions. This contrasted with previous researches about protein-like constituents existed in surface water (Yamashita and Jaffé 2008) and municipal solid waste leachate (Wu et al. 2011), in which distinct fluctuations were detected in the quenching curves with addition of Cu(II) ions. Therefore, although DOM in urban forestland soil showed similar fluorescent features to those of protein-like constituents, the intrinsic structures might be variable among different types of samples, thus affecting their quenching properties. In addition, the quenching phenomena with the addition of Ni(II) and Pb(II) ions could not be detected by traditional visual peaking method, demonstrating that EEM-PARAFAC modeling can provide complexation information at a deeper level, which might be neglected when using visual peak finding method due to prominent overlapping fluorescent components.

### Binding parameters of different types of DOM constituents

The stability constants ( $K_M$ ) and the fraction of the binding fluorescent constituents ( $f$ ) for the identified constituents and four HMs were fitted applying the Ryan and Weber model (as showed in Table 4). For Cu(II) ions, the  $\log K_M$  values were 5.31, 4.92, 5.72, and 4.90, respectively, which were in the reasonable range as those determined for agricultural soil DOM (Knoth de Zarruk et al. 2007), natural water DOM (Ryan and Weber 1982b), and lake sediments DOM samples (Xu et al. 2013). Being compared to the protein-like constituents, the higher stability constants between humic-like/fulvic-like constituents and Cu(II) ions may be related to the existence of great number of aromatic acidic functional groups (phenolic sites) (Martins and Drakenberg 1982). With respected to Ni(II) and Pb(II) ions, the binding parameters for C2 and C3 can be modeled well, whereas those for C1 and C4 cannot be modeled successfully, even if inconspicuous quenching phenomena were detected (Fig. 4a, d). These interesting results revealed that the binding mechanisms between

**Table 4** The  $\log K_M$  values for the PARAFAC constituents in urban forestland soil with four target HMs determined by Ryan and Weber model

Metals	Constituent 1			Constituent 2			Constituent 3			Constituent 4		
	$\log K_M$	$f$	$R^2$	$\log K_M$	$f$	$R^2$	$\log K_M$	$f$	$R^2$	$\log K_M$	$f$	$R^2$
Cu	5.31	67.01	0.99	4.92	50.17	0.99	5.72	63.02	0.97	4.90	70.05	0.98
Ni	FM			4.69	52.30	0.99	5.01	43.75	0.98	FM		
Pb	FM			4.37	57.01	0.92	4.69	63.89	0.98	FM		
Cd	FM				FM			FM		FM		

Ni(II)/Pb(II) ions and different types of constituents may be different. The log  $K_M$  values for Ni-C3 complex (5.01) and Pb-C3 complex (4.69) complexes were higher than Ni-C2 (4.69) and Pb-C2 (4.37) complexes, which may be attributed to the more condensed structures of C3. Furthermore, despite the EEM quenching curves of fulvic-like and humic-like constituents displayed a distinct downtrend with Cd(II) addition, the fitting results proved that the binding parameters between the four identified constituents and Cd(II) might not be analyzed reasonably through the 1:1 complex reaction model. This finding was consistent with the complexing characteristics of Cd(II) with DOM extracted from landfill leachate (Wu et al. 2011), compost-derived DOM (Xu et al. 2013), and commercial fulvic acid (Grassi and Daquino 2005).

It was noteworthy that the order of log  $K_M$  values was Cu-DOM > Ni-DOM > Pb-DOM complexes. This result can be explained as Cu(II) and Ni(II) ions may forcefully interact with DOM through the generation of comparatively stable inner-sphere complexes, where Pb(II) ions were expected to generate outer-sphere complexes to a larger extent (Hernández et al. 2006; Senesi and Loffredo 1999). In addition, the organic functional groups with strong complexing capacity usually exist in lower concentrations, whereas the organic functional groups with weak complexing capacity are present in higher concentrations (Benedetti et al. 1995; Wu and Tanoue 2001). The proportion of metal to DOM (“metal concentration effect”) is known to influence the affinity of metal-DOM. For example, Cu(II) ions incline to bind with strongest ligands (nitrogen- and sulfur-containing functional groups) of DOM at low Cu/DOM ratios (generally  $\leq 10^{-3}$ ) (Croué et al. 2003; Frenkel et al. 2000). Moreover, according to the Hard-Soft-Acid-Base theory (Haitzer et al. 2002), this effect was believed to be powerful for Ni(II) and Pb(II) due to their “hard-medial-soft” property and the low content of “soft” ligands compared to the high concentration of “hard” oxygen ligands in DOM. In this study, all the minimum initial values of Cu/DOM, Ni/DOM, and Pb/DOM were approximately 0.05, resulting in a rapid saturation of the powerful binding sites and that the most of the above three HMs tended to bind the weak sites, such as hydroxyl, phenolic, carboxylic, and carbonyl. Therefore, the log  $K_M$  values fitted by the quenching curves should be regarded as theoretical averaged constants.

FM: failed to be modeled.

### Risk assessment of soil environments

EEM-PARAFAC had been shown to discriminate urban forestland soil-derived DOM into various significative components and evaluate the relative abundances of these different organic substances. Since PARAFAC modeling has the ability to determine the binding parameters of the identified organic constituents with HMs, it is possible to assess the complexation information of metal-DOM speciation in greater details.

Moreover, as diffusely recognized, DOM plays a vital role in the biotoxicity and transportation of HMs in soil environments (Huang et al. 2019; Kalbitz and Wennrich 1998; Uchimiya et al. 2010; Weng et al. 2002). A relatively stronger binding ability in the existence of DOM indicated that urban forestland soil played a more significant role in the detoxification of HMs. The toxicity of HMs may be weakened after complexing with different types of ligands (Sheng et al. 2013). However, recent researchers demonstrated that elevated HM concentrations in soil could be attributed to stormwater runoff, groundwater flow, atmospheric cycling, vehicle exhaust, waste disposal, and industrial emission (Hou et al. 2017). Therefore, based on the strong metal-DOM affinity found in this study, it was deduced that the contamination of HMs in urban forestland soil might become more terrible due to inadequate control of pollution sources and neglect of maintenance for soil quality.

In addition, the content level and the chemical composition of DOM should be considered as the dominating factors in evaluating the risk of HMs in urban forestland soil environments. Previous studies on risk evaluation of HMs in soil were primarily concentrated on the total concentration and speciation, applying multiple statistical methods (principal component analysis, cluster analysis, Pearson correlation analysis, etc.) to identify the soil properties (Ali et al. 2016; Huang et al. 2015; Kelepertzis 2014). The various quenching behaviors of Cu(II), Cd(II), Pb(II), and Ni(II) found in this study suggested that the strength of complexing interaction should also be regarded as the risk assessment of HMs. Furthermore, EEM-PARAFAC model and titration results revealed that the highest binding affinity was Cu-DOM, followed by Ni-DOM, and the worst was Pb-DOM (the log  $K_M$  values of Cd-DOM cannot be obtained). Therefore, Pb(II) or Cd(II) may present a stronger toxicity in urban forestland soil ecosystems due to its relative low binding capacities and thus their toxicity should not be ignored for engineering and scientific researchers.

### Conclusions

In this study, urban forestland soil samples in Daxing District were scientifically sampled with the help of GIS technology. The UV-vis and fluorescence characteristics showed that the DOM in urban forestland soil contained a complex mixture of terrestrial and microbially derived organic substances with smaller molecular weight and more aromatic structures. EEM and PARAFAC analyses were combined to study the binding affinity of HMs to urban forestland soil DOM. Four independent fluorescent constituents (1 fulvic-like, 2 humic-like, and 1 proteins-like) were decomposed by the PARAFAC model. Being compared to protein-like constituents, the binding ability of fulvic-/humic-like constituents with Cu(II) ions were relatively stronger. Considering that the Ryan and Weber



model only successfully calculated the binding parameters of Ni(II)/Pb(II) ions and humic-like constituents thus indicate that their corresponding complexes exhibited higher stability constants than the fulvic-like and protein-like constituents. Recognizing the vital role that DOM in urban forestland soil can influence the mobility, bioavailability, and toxicity of HMs, plentiful strategies should be proposed for the management of urban forestland soil microenvironments. The above results provided a detailed theoretical basis of the binding affinities of HMs and the different types of DOM at the molecular level, which was pivotal for efficient control of the transportation and transformation of HMs in urban forestland soil.

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